

# Isotopic techniques for tracing evapotranspiration at the field scale. Simple or complex?

**David G. Williams**

University of Wyoming  
Laramie, Wyoming, USA

**Aaron Van Pelt**

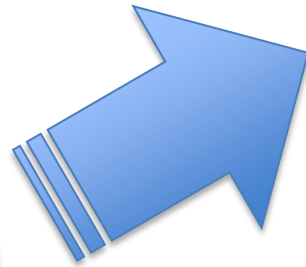
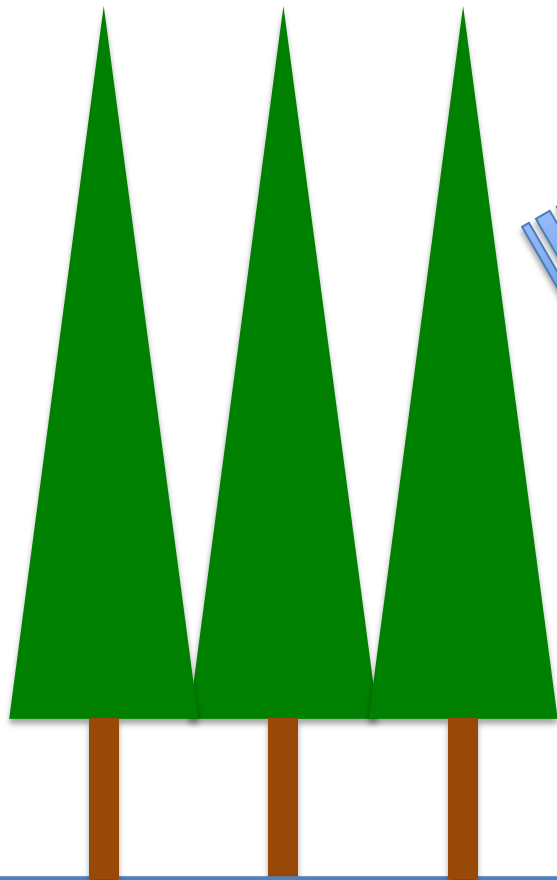
Picarro, Inc.  
Sunnyvale, California, USA

**Leo Mayr and Lee Heng**

FAO/IAEA Agriculture and  
Biotechnology Laboratory,  
Seibersdorf, Austria

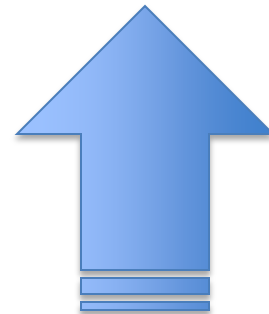


# Isotopes of H and O in atmospheric water vapor reveal sources of evapotranspiration



Plant transpiration  
 $\delta^2\text{H}$  and  $\delta^{18}\text{O} = ???$

Soil evaporation  
 $\delta^2\text{H}$  and  $\delta^{18}\text{O} = ???$

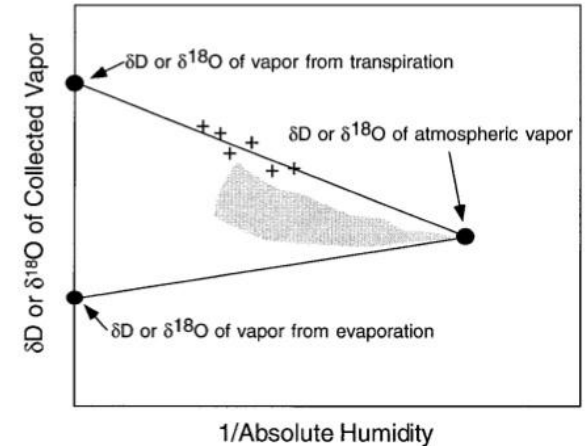


Soil water  $\delta^2\text{H}$  and  $\delta^{18}\text{O}$

# Isotope approaches for separating ET fluxes

## Isotope mass balance - Keeling plots

*Relatively simple*



## Isotope flux gradient approach

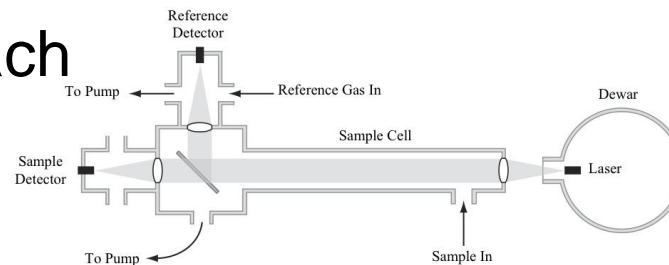
*Somewhat complicated*

centre of the fields. Significantly, these were distinct from the short-re calculated based on leaf scale irrently with the air sampling. As  $\delta_p^{13}$  were consistent with  $\delta_b^{13}$  values lyses (Table 1). Note that  $\delta_p^{13}$  and 0‰ higher in the corn field than ove C<sub>4</sub> plants. The much smaller eat and corn fields reflected small ment and stomatal conductance

$$T = -\bar{K} \times \frac{\Delta C_T}{\Delta_z} = f_1 \left( -\bar{K} \times \frac{C_b}{\Delta_z} \right) = f_1 \times ET$$

## Eddy covariance approach

*Very complicated  
...and expensive*

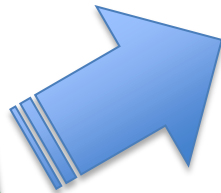
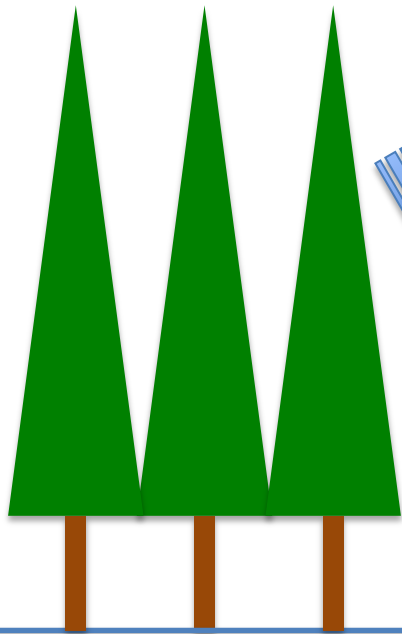


$$F^x = \bar{\rho}_a \overline{w'x_w^x} + S^x$$

$$d_F = \frac{\frac{\partial F^{18}}{\partial t}}{\frac{\partial F^{16}}{\partial t}} - \frac{1}{R_{std}} \cdot 1000$$

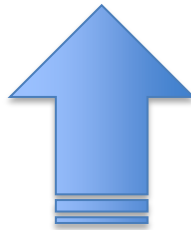
# Assumptions of Keeling plot approach

- Source and background values are stable
- Only evapotranspiration (no dew formation)
- Isotopic steady state of plant transpiration

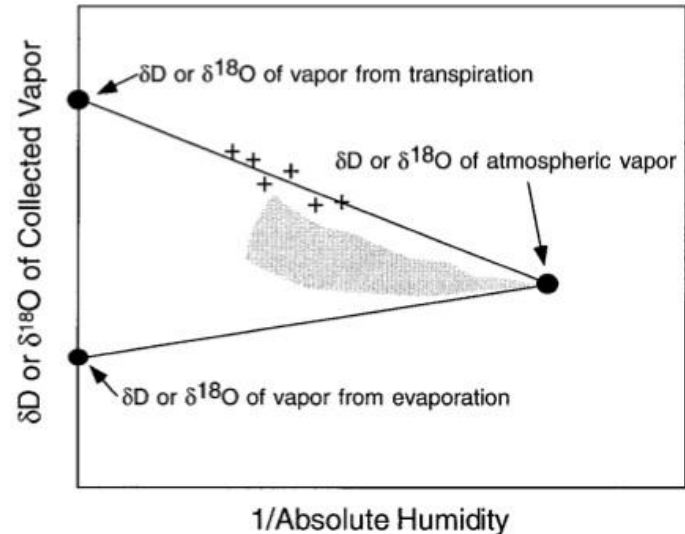


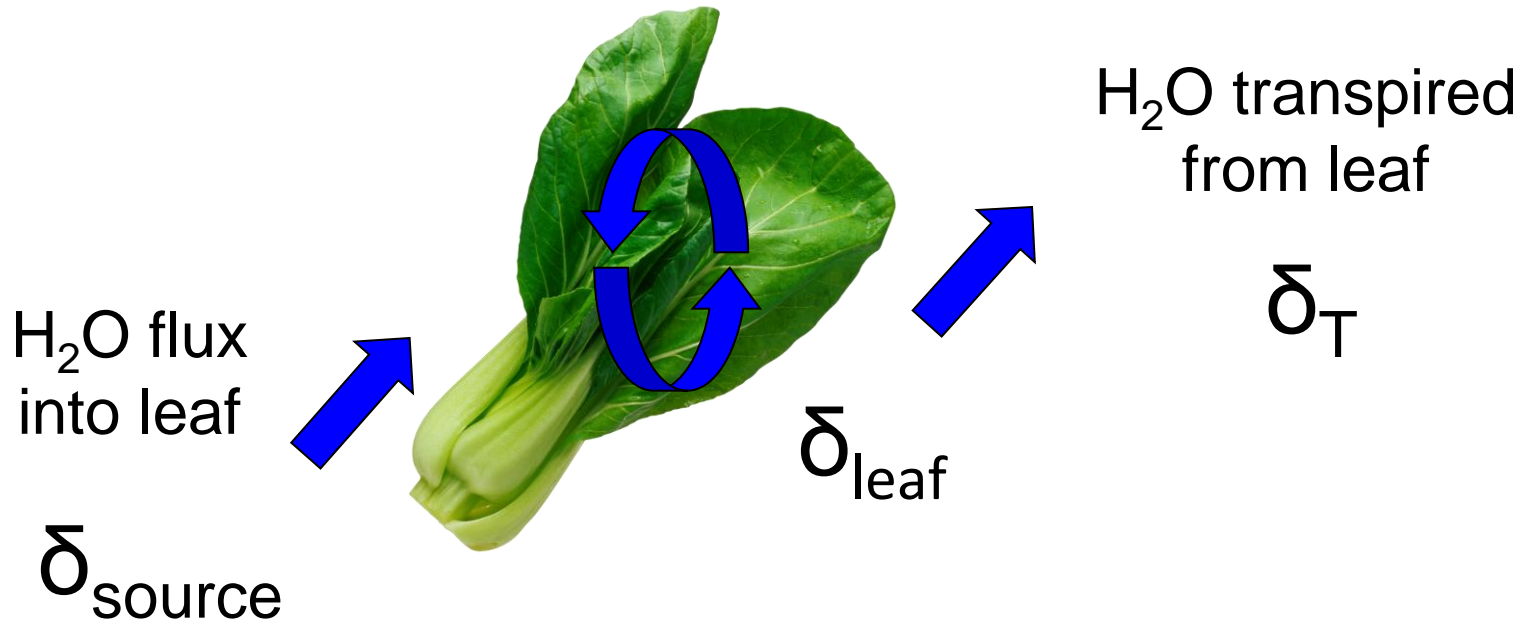
Plant transpiration  
 $\delta^2\text{H}$  and  $\delta^{18}\text{O} = ???$

Soil evaporation  
 $\delta^2\text{H}$  and  $\delta^{18}\text{O} = ???$



Soil water  $\delta^2\text{H}$  and  $\delta^{18}\text{O}$

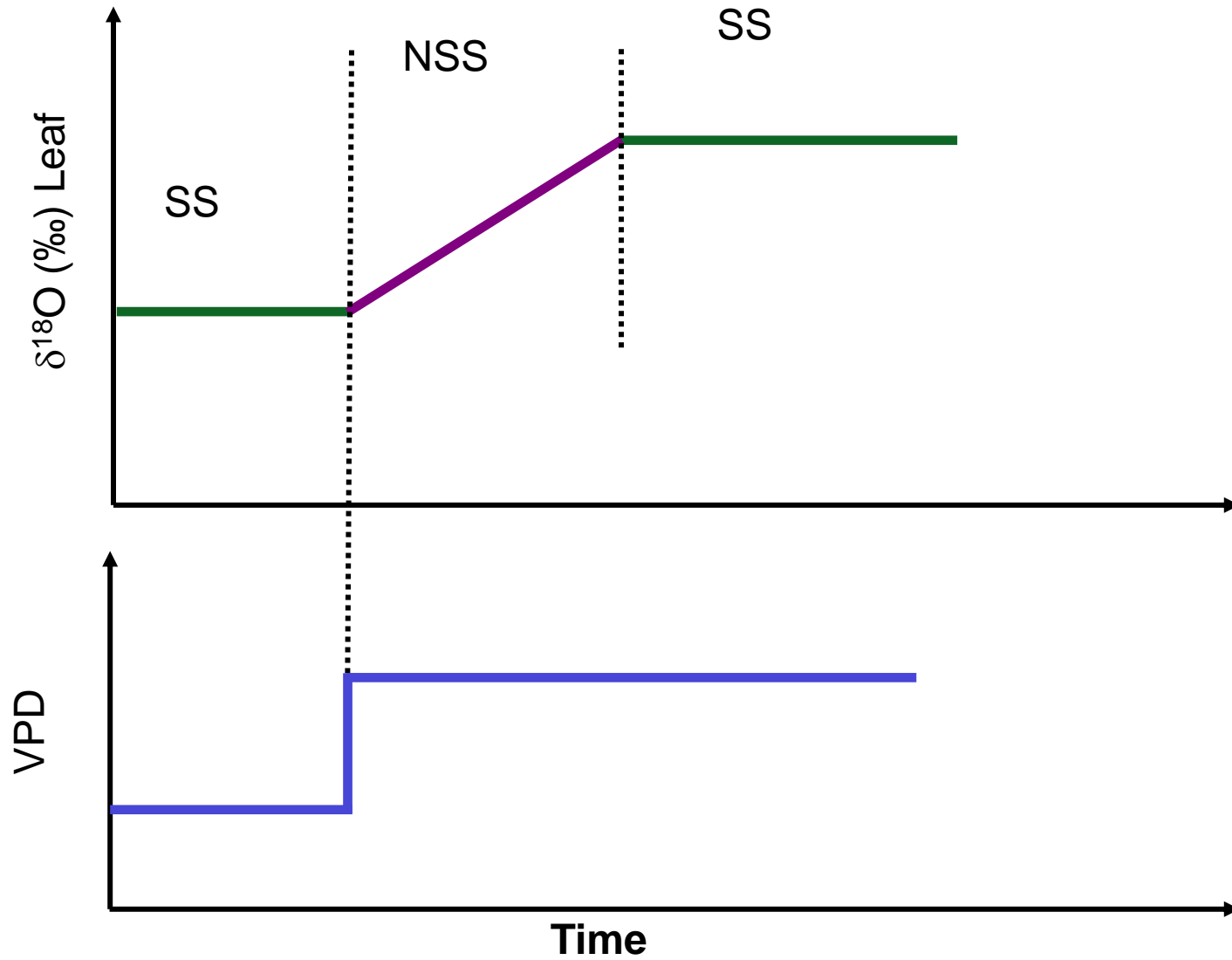




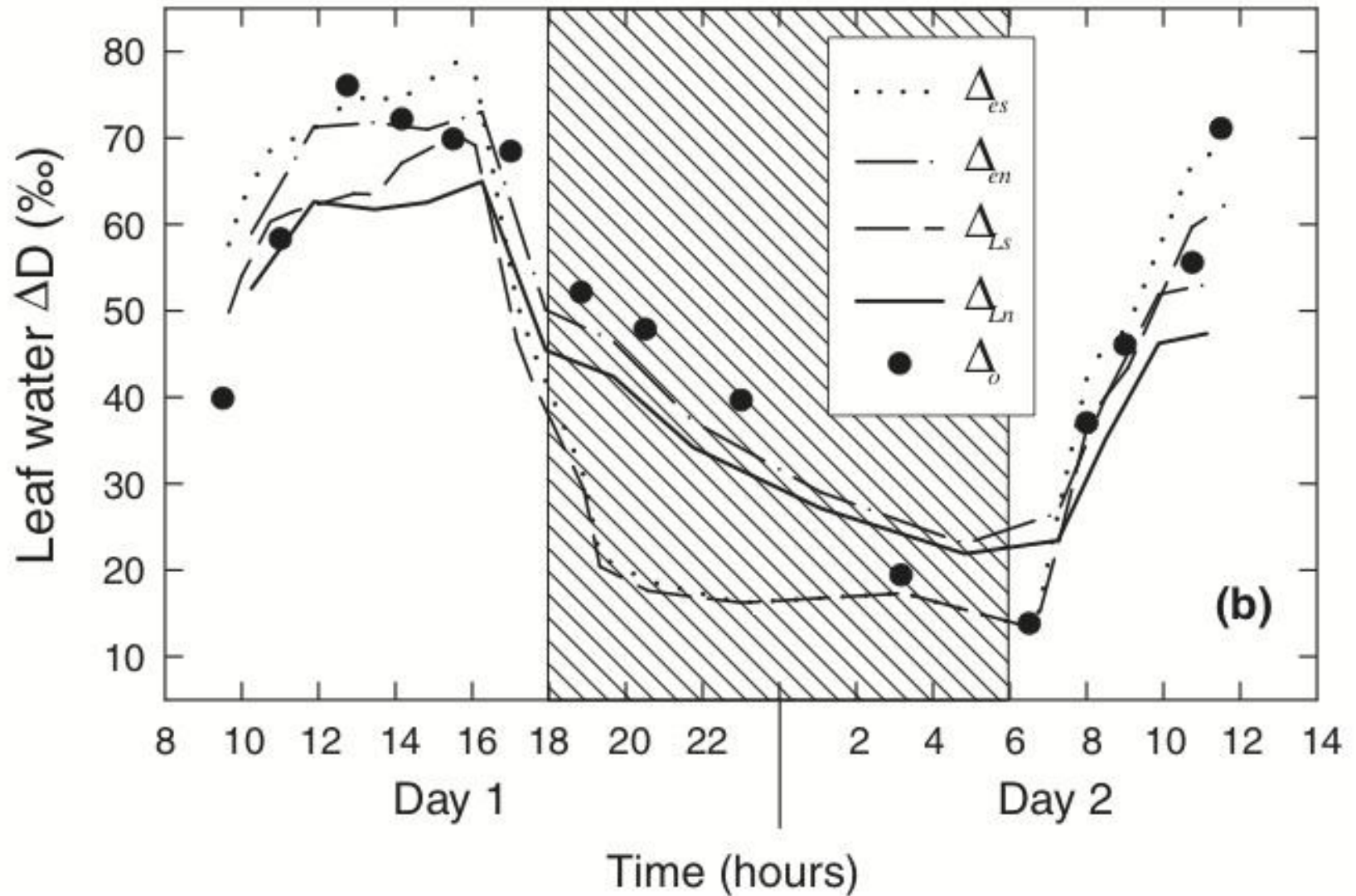
At isotopic steady state,  $\delta_{\text{T}} = \delta_{\text{source}}$

At isotopic non-steady state,  $\delta_{\text{T}} \neq \delta_{\text{source}}$

# Non steady state



# Leaf transpiration at isotopic steady state



# Glacier Lakes Ecosystem Experiments Site (GLEES) U.S. Forest Service



2008/09/27



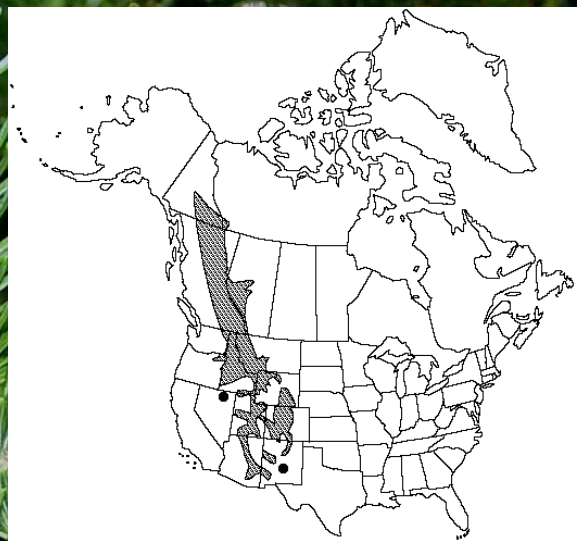
0 yr



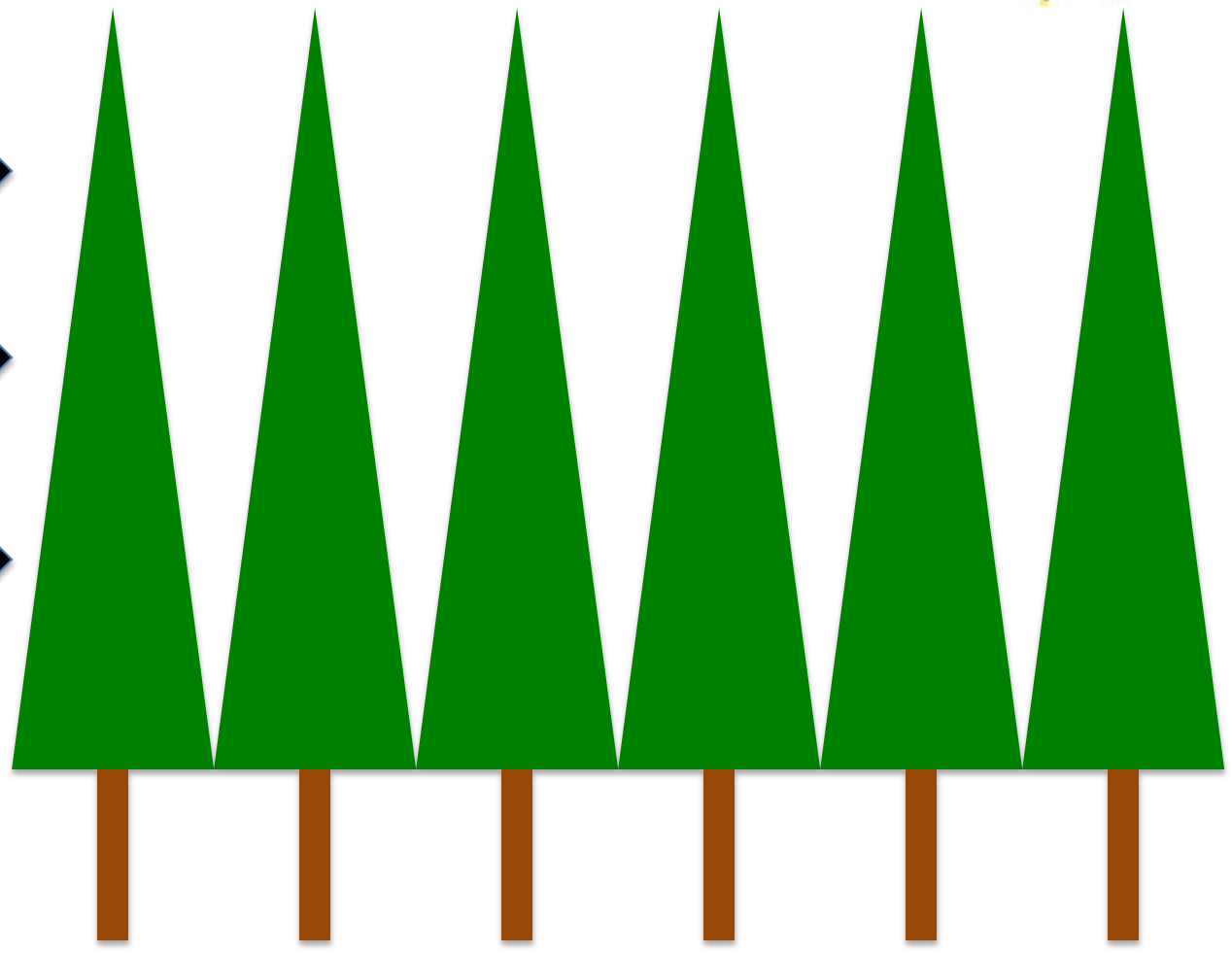
1 yr



3 yr



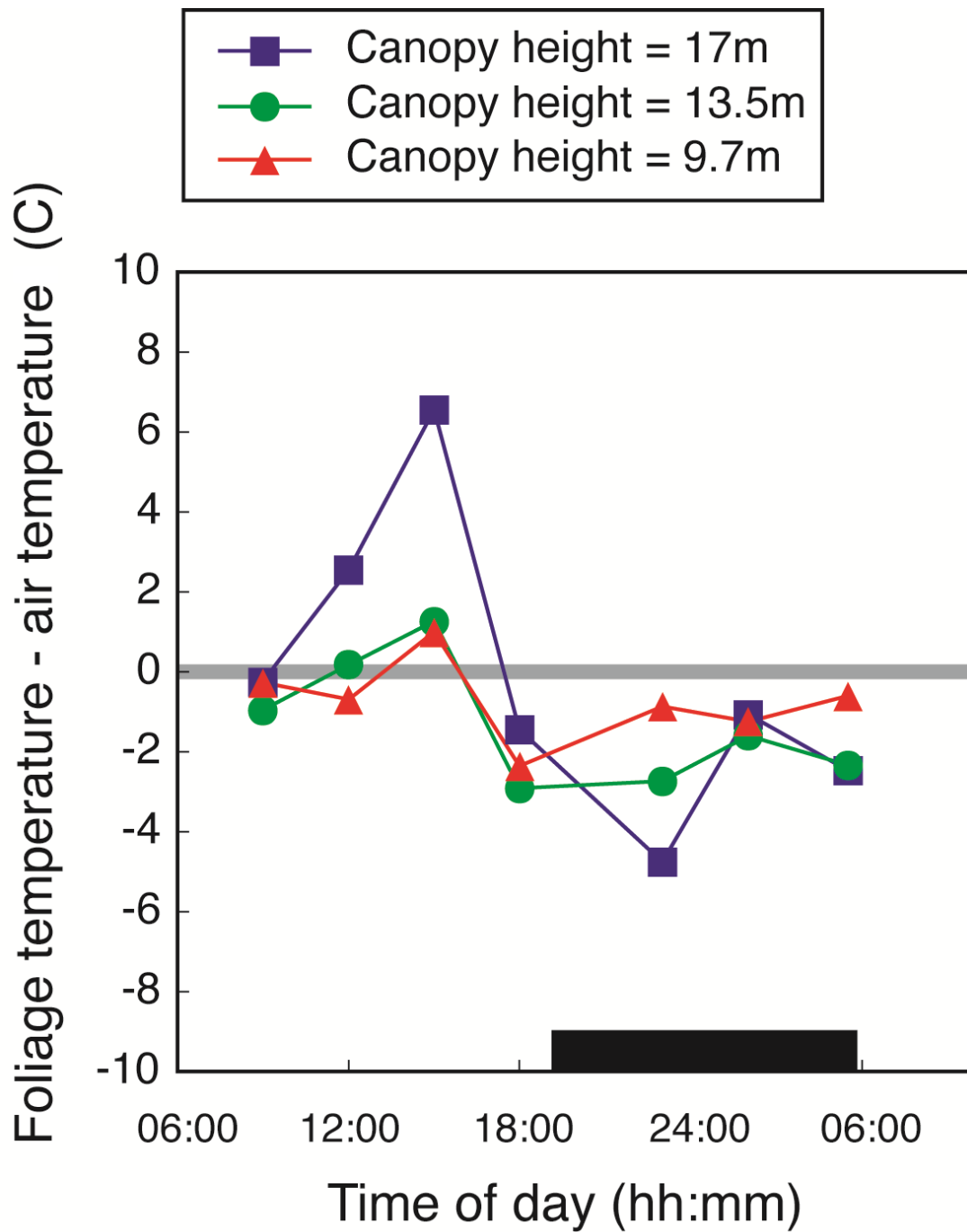
*Abies lasiocarpa*

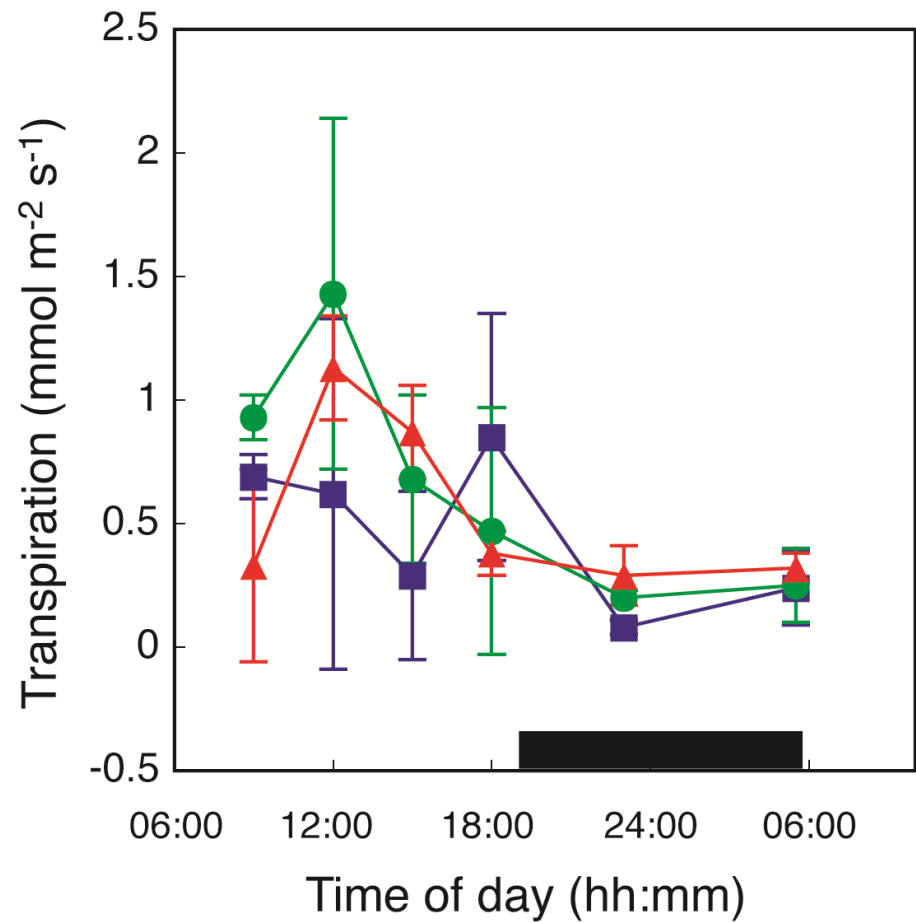
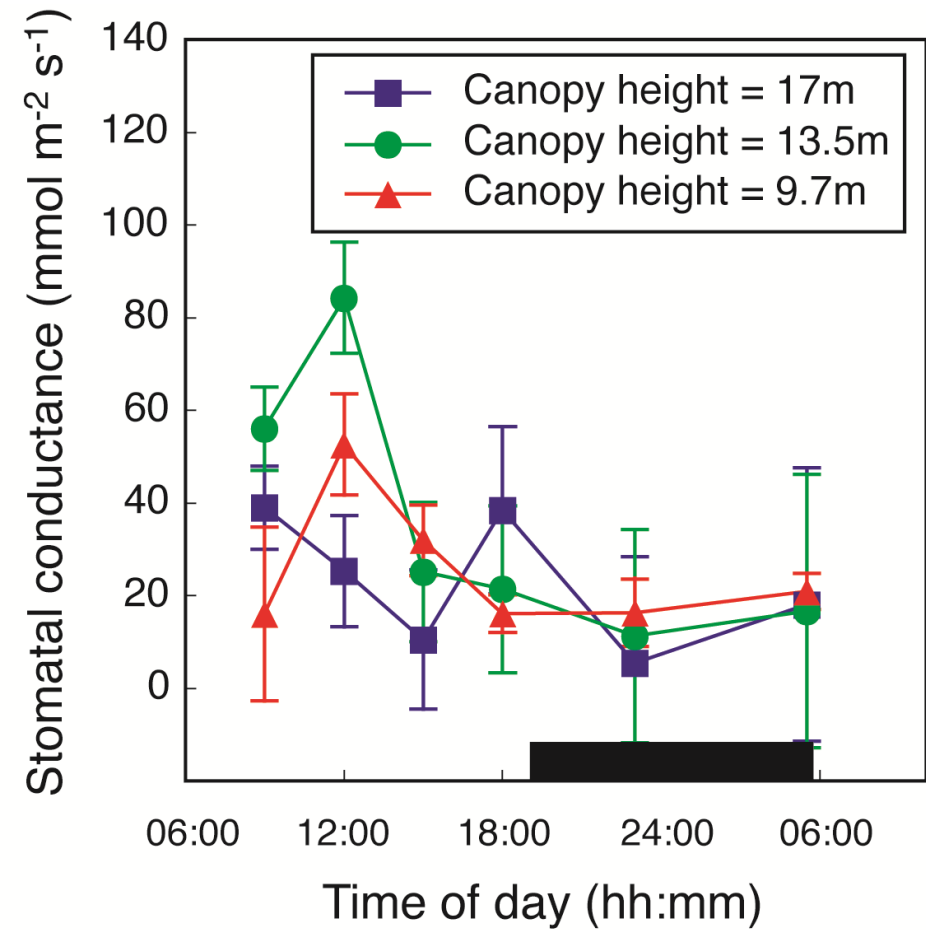


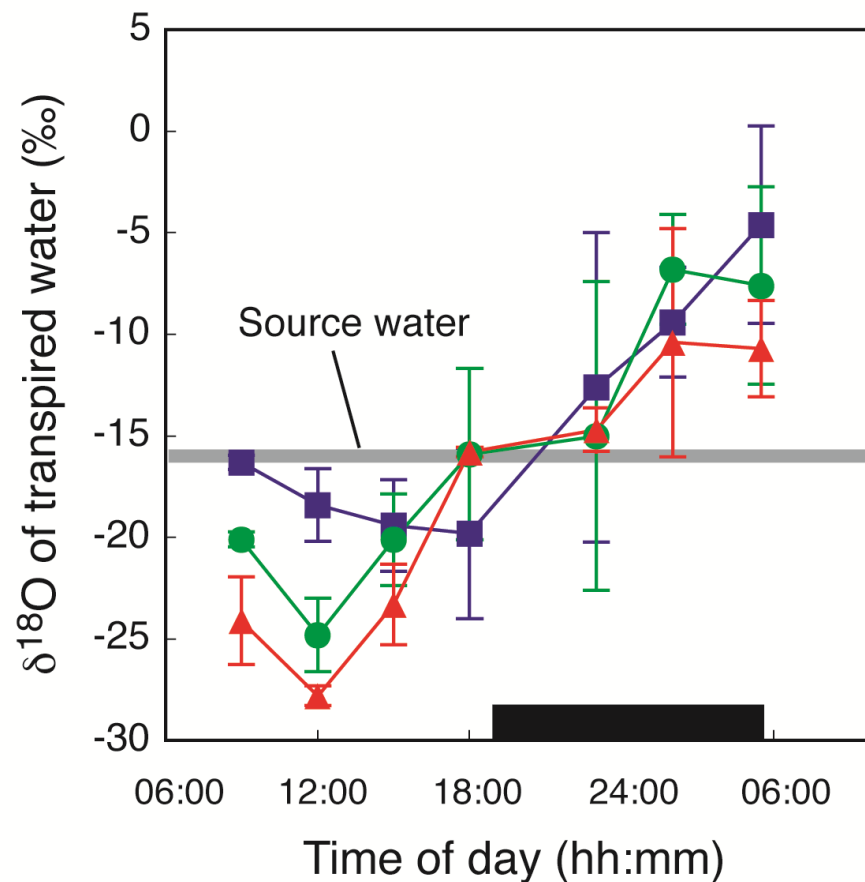
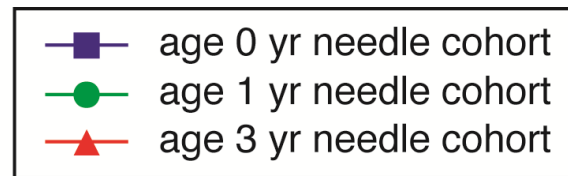
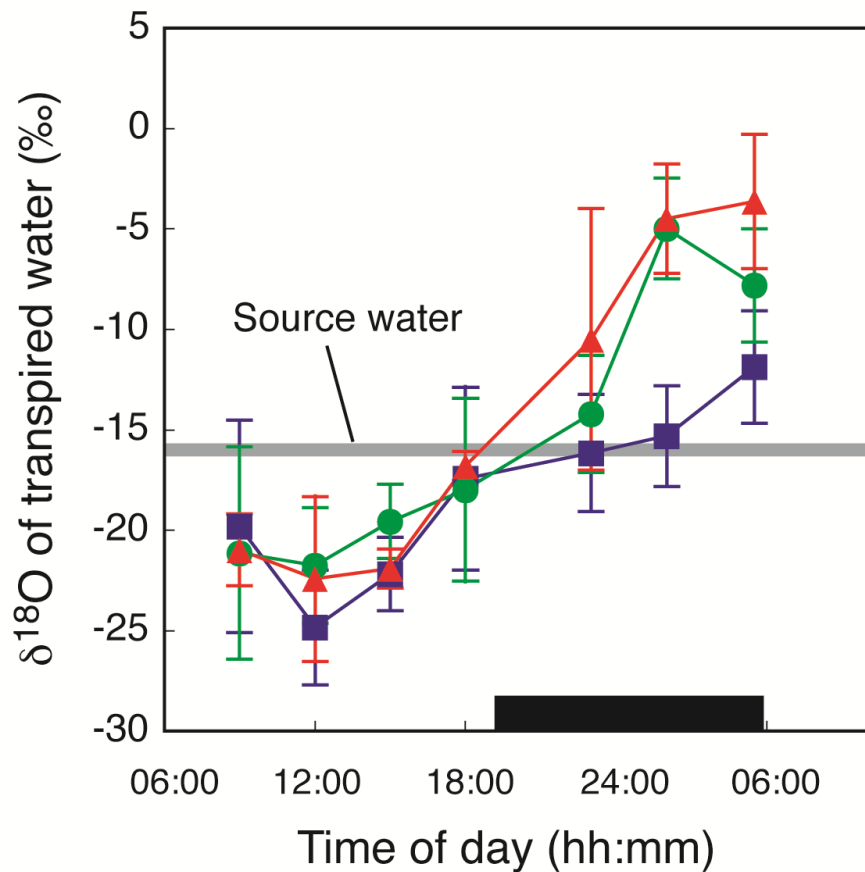
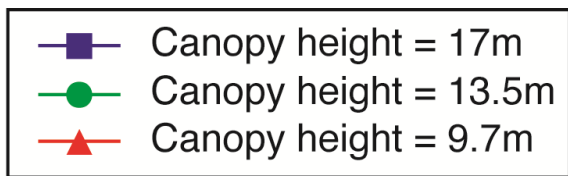
17m →

13.5m →

9.7m →









Transpiration in *Abies lasiocarpa* operated predominantly at isotopic **non-steady state**

Systematic variation in leaf water  $^{18}\text{O}$  enrichment and transpired water with canopy height and leaf age

Next step:  
Model isotopic non-steady state of canopy transpiration considering complexity of canopy physiological processes and micro-environment